

Aggregation and Uncertainty

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with

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Level-2 versus Level-3

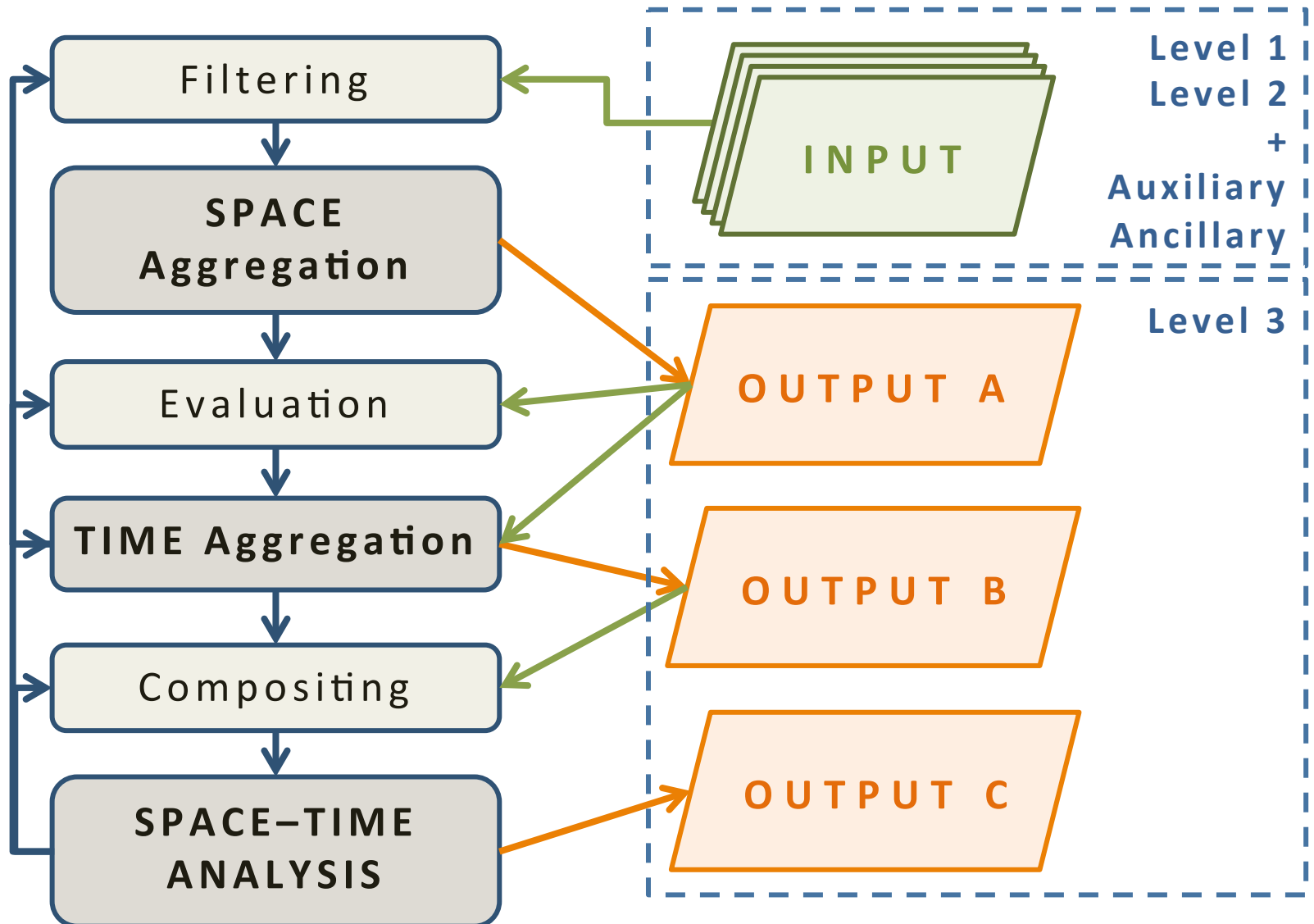
In both cases, y is not measured directly, but is determined from n other quantities x_1, x_2, \dots, x_n through a functional relationship f :

$$y = f(x_1, x_2, \dots, x_n)$$

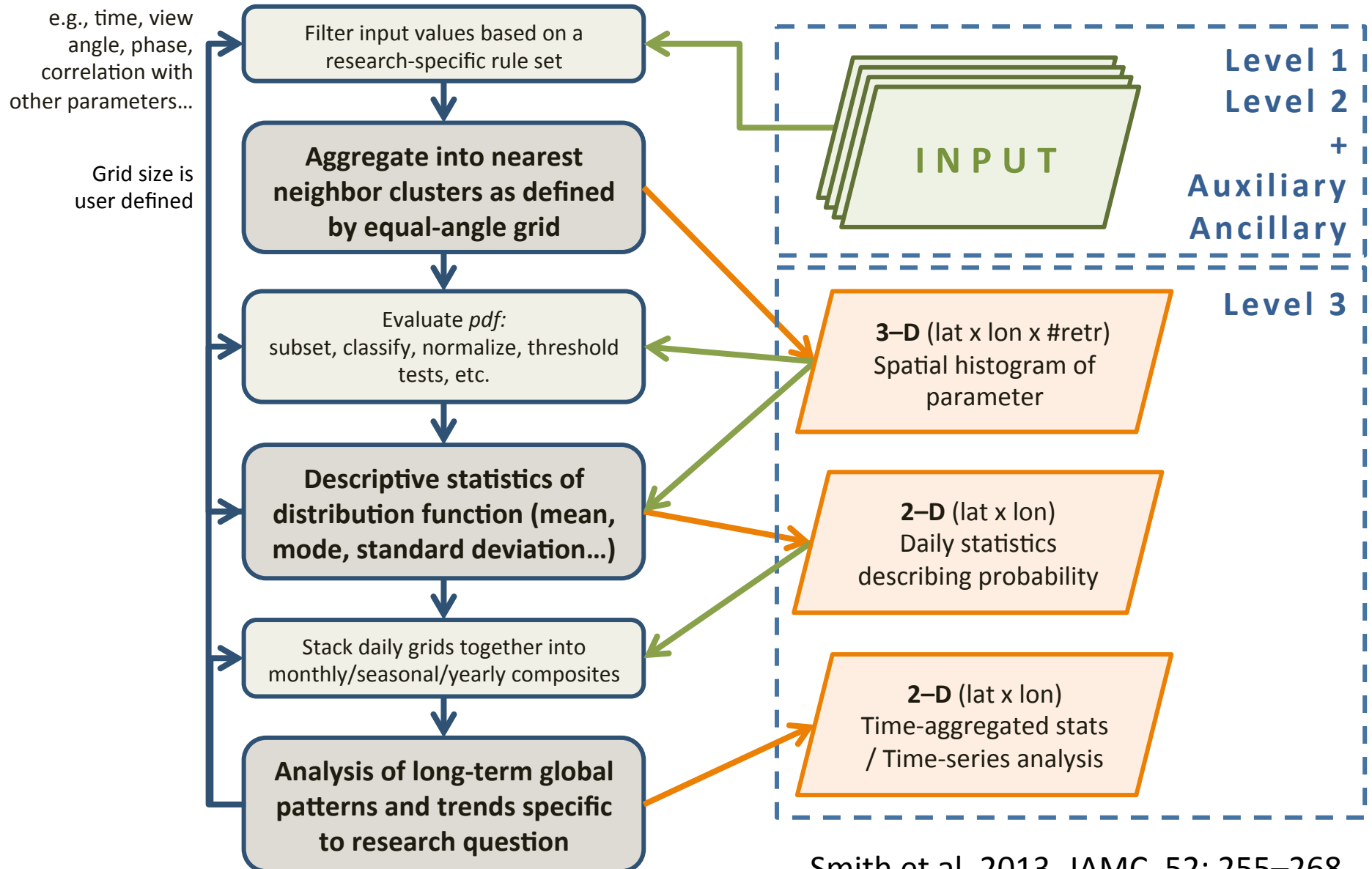
Level 2	Level 3
y = atmospheric parameter retrieval (weather observation)	y = atmospheric parameter aggregate (climate observation)
f = retrieval/inversion algorithm	f = aggregation algorithm
x = top of atmosphere measured spectrum. Number of spectral bands/channels determines magnitude of n	x = L2 retrieval parameters. Both L2 pixel size and L3 grid size determines magnitude of n
Sources of uncertainty include forward model, a priori, instrument noise, spectral resolution, inversion algorithm, etc.	Sources of uncertainty include Level 2 retrieval uncertainty, aggregation algorithm, scene interpretation, etc.
Inversion problem ranges from near-linear to non-linear	Aggregation problem range from near-linear to non-linear
Solution is best estimate of probability distribution; a mean with SD. Both input and output are <i>probabilities</i>	Solution is best estimate of probability distribution; a mean with SD. Both input and output are <i>probabilities</i>

UW Space–Time–Gridding (STG) Framework

Modular design allows ease of processing/experimentation and transparency of data flow



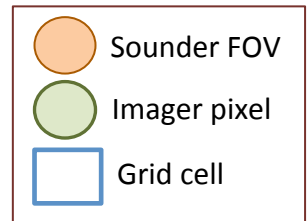
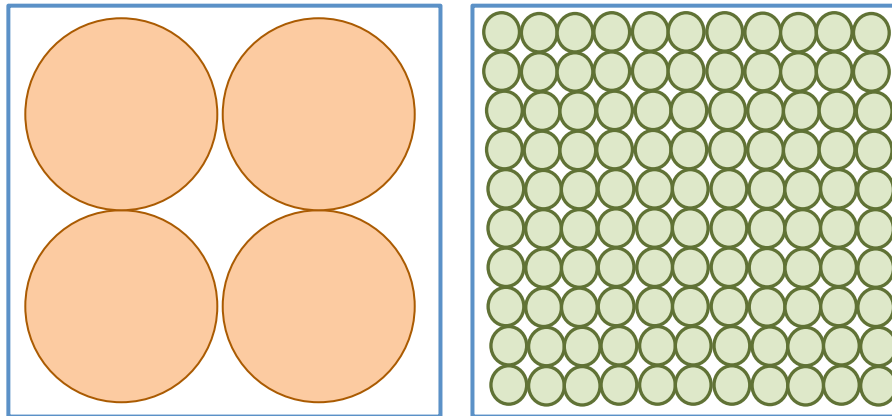
UW Space–Time–Gridding (STG) Framework



Aggregation as a global co-location scheme

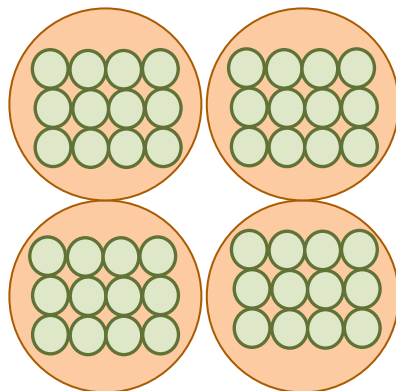
A study in Brightness Temperature (BT): comparing Imagers (VIIRS) and Sounders (CrIS)

STG method



- Convolve Sounder BT to Imager BT
- Aggregate Sounder BT (avg + stdev)
- Aggregate Imager BT (avg + stdev)
- Select uniform scenes:
 - (1) Sounder stdev < threshold 1
 - (2) Imager stdev < threshold 2

Traditional pixel-by-pixel colocation method



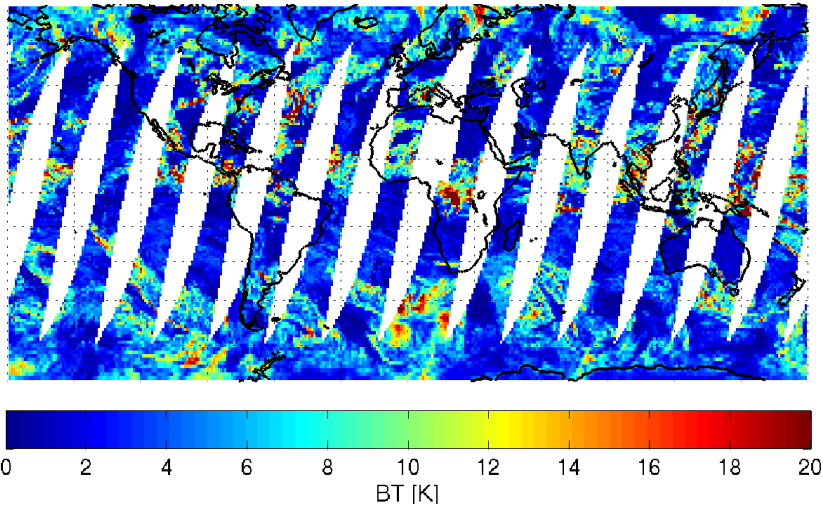
- Convolve Sounder BT to Imager BT
- Colocate Imager BT to sounder FOV (avg + stdev)
- Select uniform scenes:
 - (1) Imager stdev < threshold 2

Bias = Average difference of uniform scenes

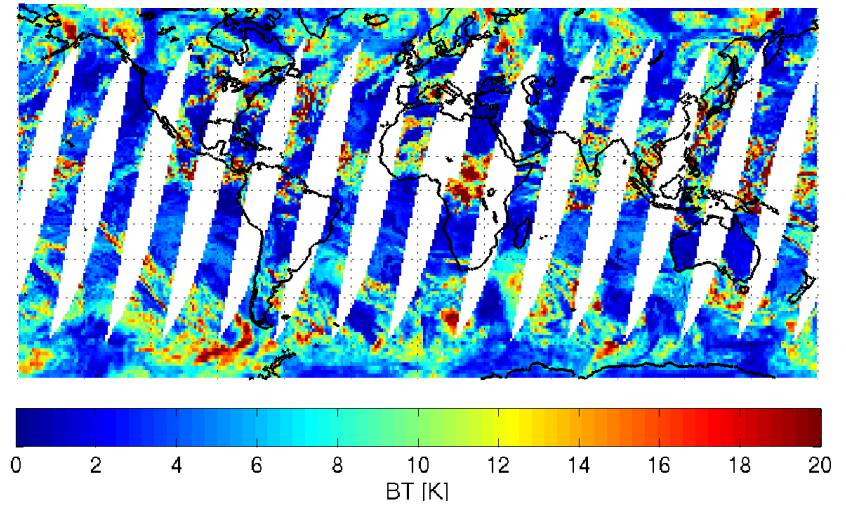
Error = Standard deviation of difference of uniform scenes

Aggregation as a global co-location scheme

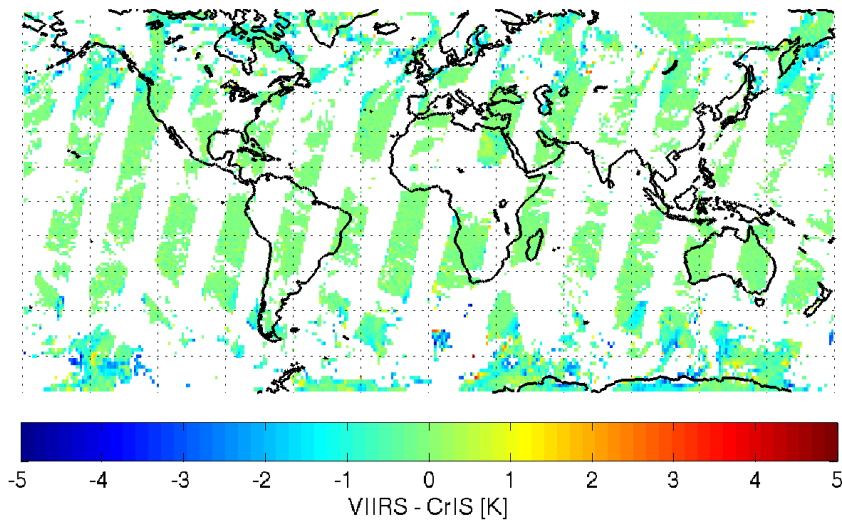
Convolved CrIS BT @ 10.8 μm 1-degree SD



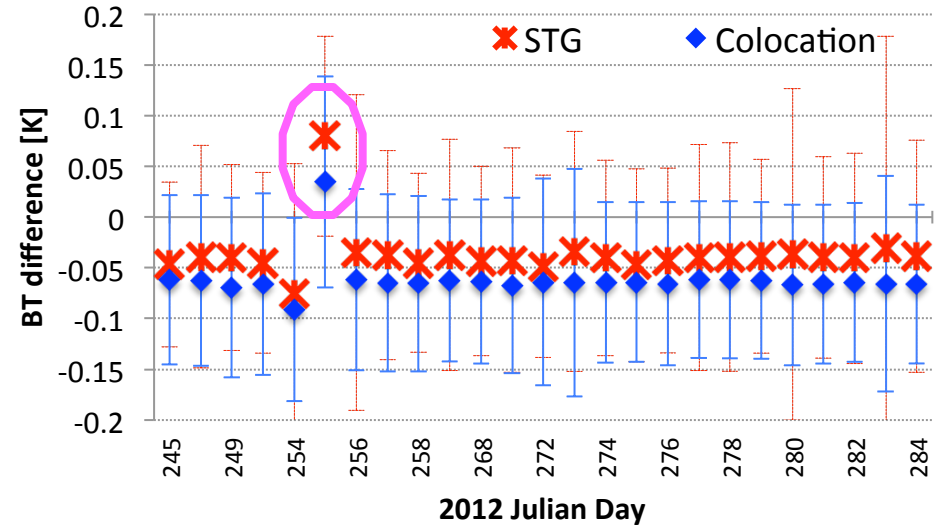
VIIRS BT @ 10.8 μm 1-degree SD



For uniform scenes: VIIRS SD minus CrIS SD



Global average of uniform-scene-differences



For global/zonal statistics, if aggregation is done correctly, co-location error is minimal

Applications

Propagation of L2 uncertainty

Instrument:

MODIS

Retrieval Algorithm:

Coll.5 (C5) & Coll.6 (C6)

Parameter:

**Cloud Effective Radius
[μm]**

- Daytime only (solar zenith angle)
- Water clouds (cloud phase retrieval)
- Retrieval Quality/Uncertainty flags

1-degree equal-angle grid

Sample size threshold

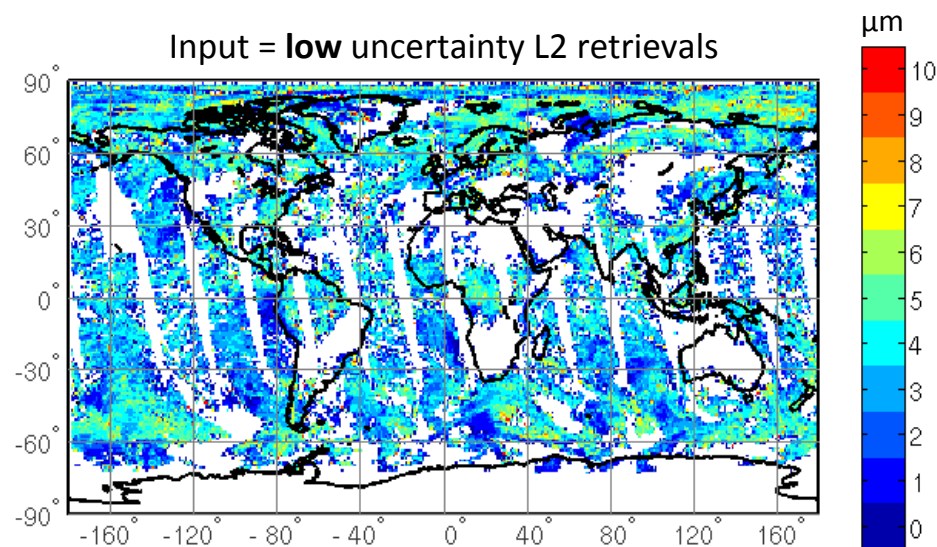
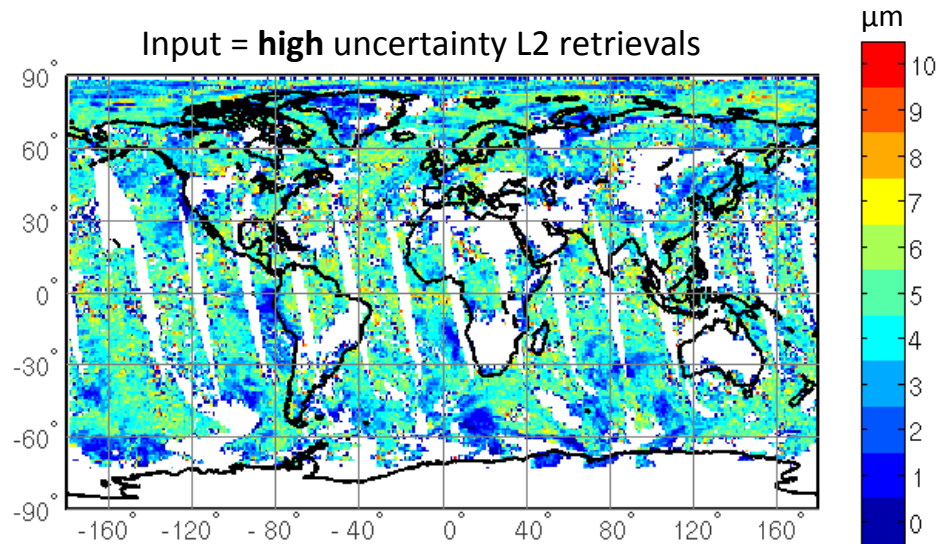
Arithmetic mean + SD

Monthly stats: September 2012

Can the propagation of uncertainty be measured (L2 to L3)? Can uncertainty in L3 be reduced by applying appropriate L2 filters?

Propagation of L2 uncertainty

Standard deviation (SD) of retrievals on 1-degree grid



Does L2 uncertainty affect L3 uncertainty?

High uncertainty L2 = using all retrievals irrespective of value in retrieval uncertainty flag

Low uncertainty L2 = using **only** those retrievals with < 10% in uncertainty flag

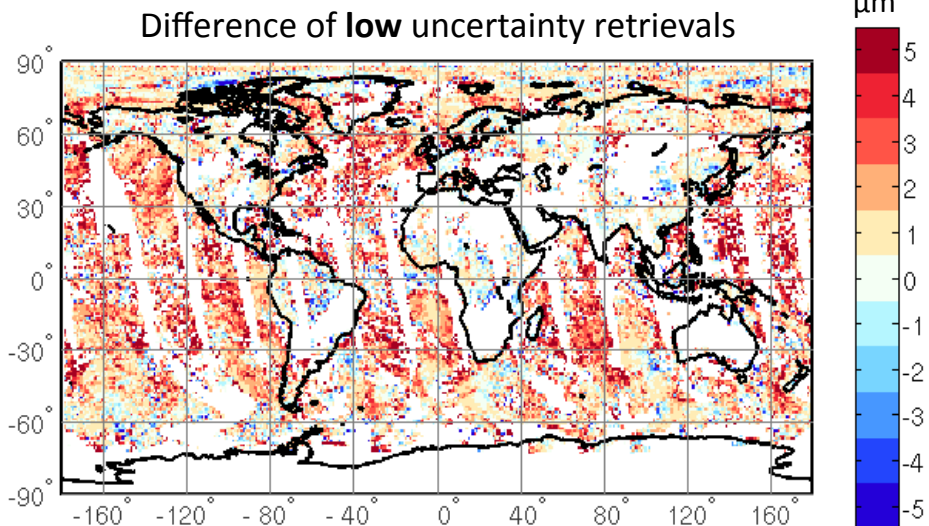
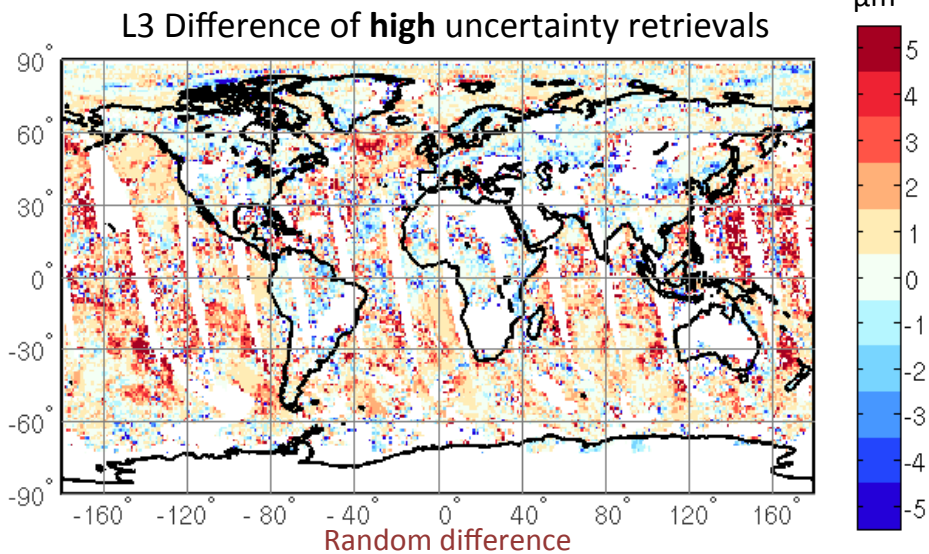
L2 uncertainty is a retrieval by-product

L3 uncertainty is measured as the SD of gridded probability distribution (sample of retrievals per grid cell)

- L3 uncertainty is significantly reduced across the globe when L2 uncertainty is reduced.
- More stringent filtering rules result in less coverage but more reliable results

Propagation of L2 uncertainty

Daily average of retrievals on 1-degree grid



How does L2 uncertainty affect product comparisons?

$L3 \text{ Difference} = C5 \text{ minus } C6$

High uncertainty L2 = using all retrievals irrespective of retrieval quality

Low uncertainty L2 = using **only** those retrievals with < 10% uncertainty

L2 uncertainty is a retrieval by-product

L3 uncertainty is measured as the SD of gridded probability distribution (sample of retrievals per grid cell)

- L2 filtering strongly affects differences among data products
- Using only low uncertainty L2 retrievals in L3, highlights systematic differences between two retrieval algorithms

Characterizing uncertainty: random vs. systematic

Instrument:

MODIS

Retrieval Algorithm:

C5 & C6

Parameter:

**Cloud Top Pressure
(CTP, [hPa])**

Difference:

C5 minus C6

- Daytime only (solar zenith angle)
- Near-nadir (sensor zenith angle)
- Retrieval Quality flags

1-degree equal-angle grid

Sample size threshold
Classify into ordinal classes: low/mid/high

Arithmetic mean + SD

Monthly stats: September 2012

**Can aggregation be used to identify
systematic differences in L2 products?
How much of the observed difference is
random?**

Systematic Differences between two Algorithms

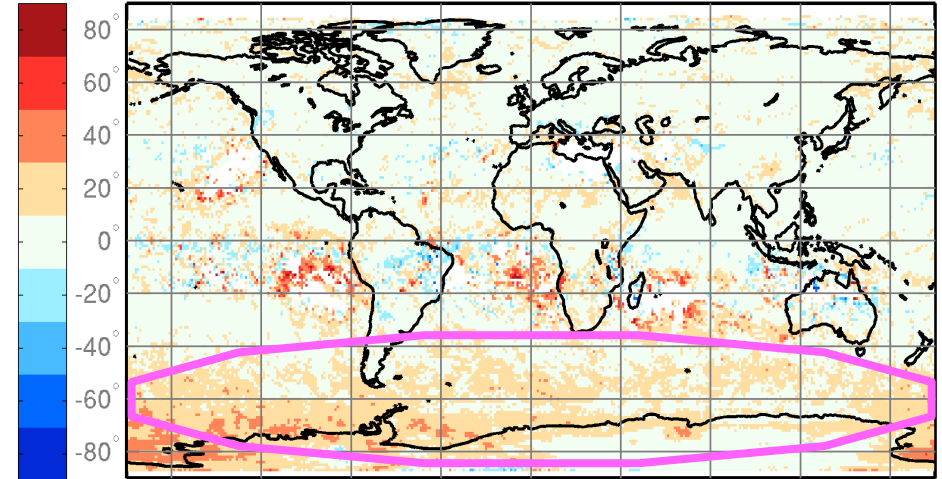
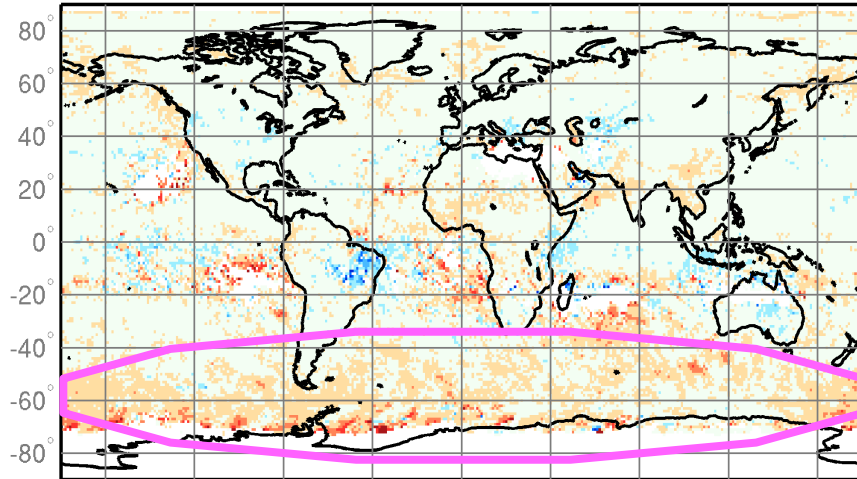
Day Time

High clouds < 440 hPa

Night Time

MODIS C5 minus C6 avg difference Aug 2012 - Daytime high clouds (9.63 hPa)

[hPa] MODIS C5 minus C6 avg difference Aug 2012 - Night time high clouds (13.34 hPa)



C6 has higher cirrus clouds, day and night

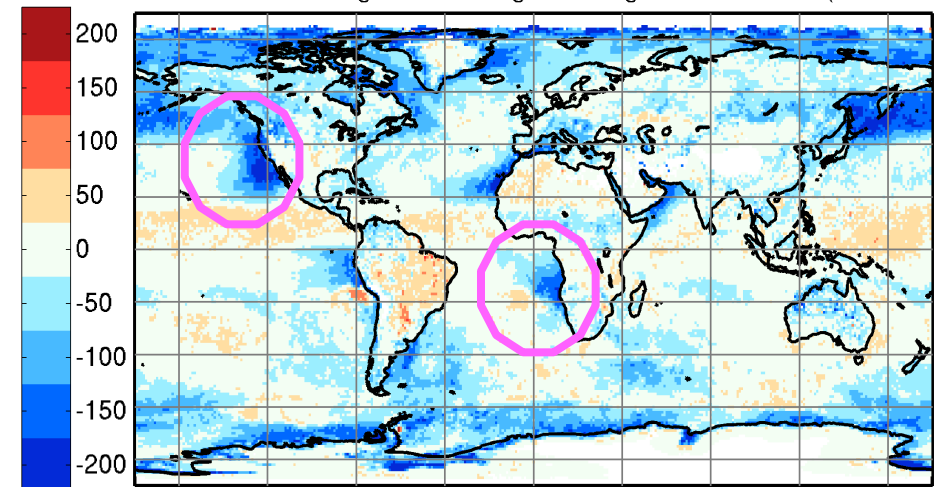
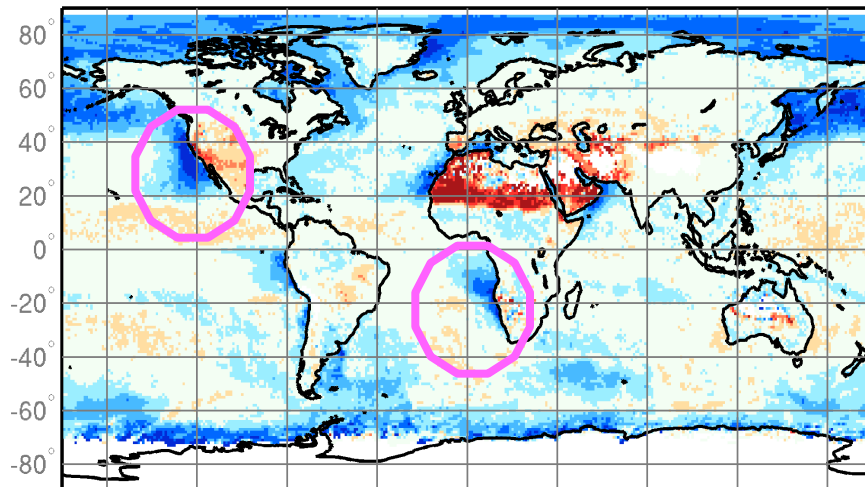
Day Time

Low clouds > 680 hPa

Night Time

MODIS C5 minus C6 avg difference Aug 2012 - Daytime low clouds (-13.44 hPa)

[hPa] MODIS C5 minus C6 avg difference Aug 2012 - Night time low clouds (-20.61 hPa)



C6 has lower marine stratocumulus clouds, day and night

Systematic Differences between two Algorithms

Summer - Average

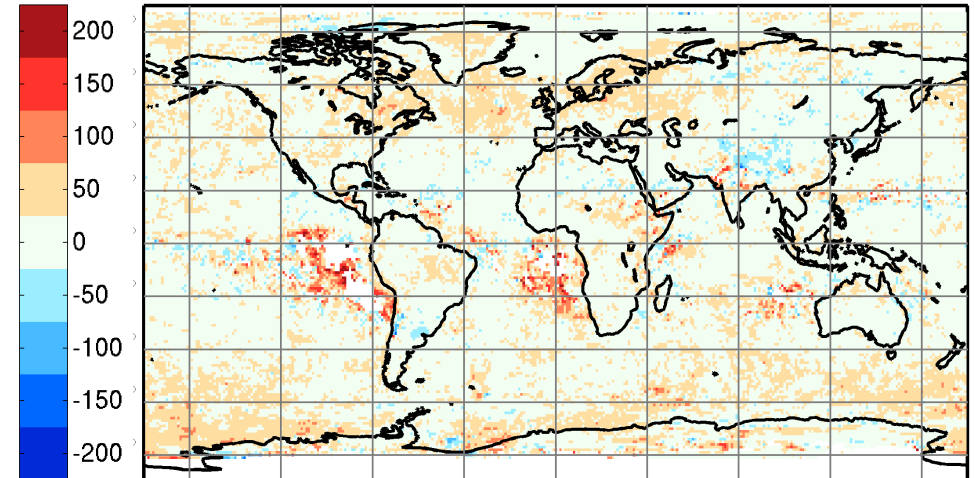
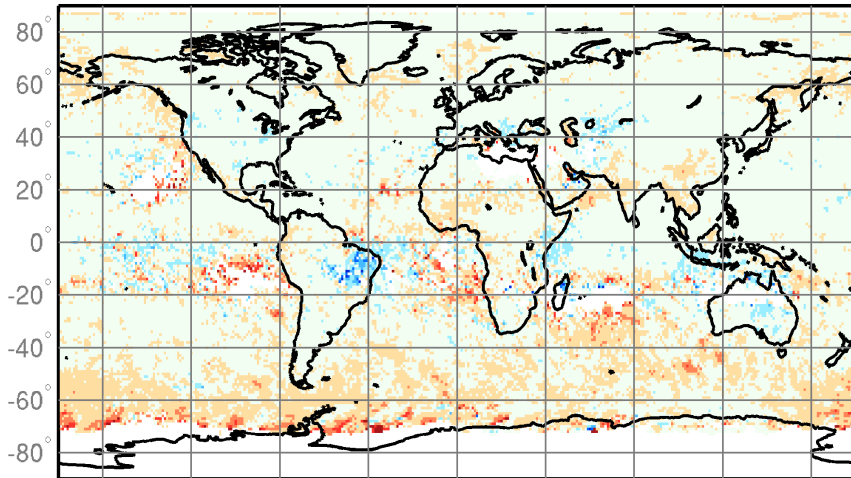
High clouds < 440 hPa

Winter - Average

MODIS C5 minus C6 avg difference Aug 2012 - Daytime high clouds (9.63 hPa)

[hPa]

MODIS C5 minus C6 avg difference Nov 2012 - Daytime high clouds (12.45 hPa)



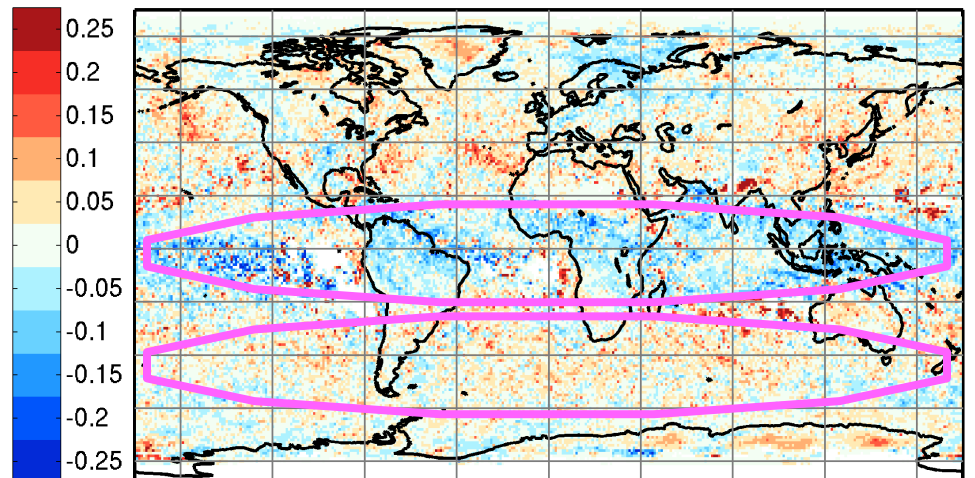
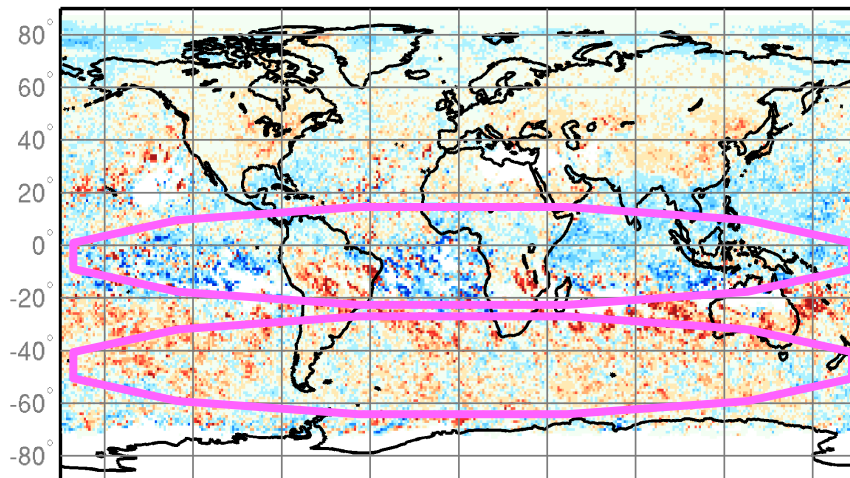
Summer - Frequency

Winter - Frequency

MODIS C5 minus C6 perc difference Aug 2012 - Day time high clouds (NaN)

[%]

MODIS C5 minus C6 perc difference Nov 2012 - Day time high clouds (NaN)

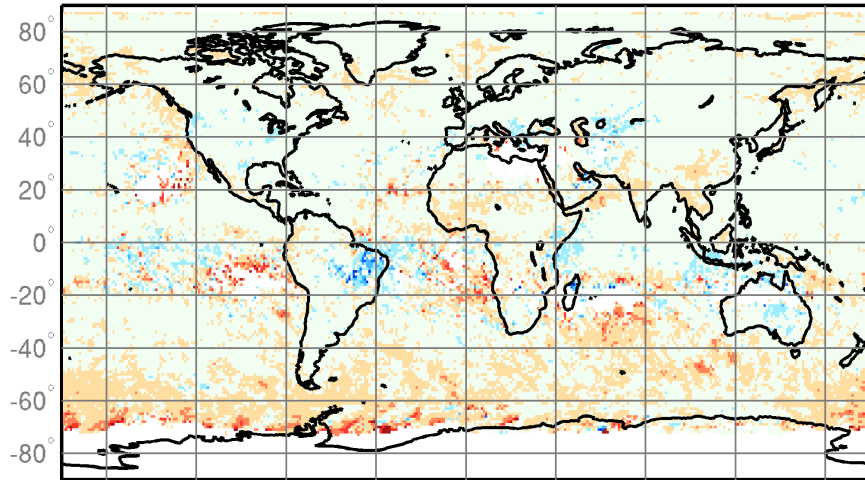


C6 frequency of high clouds higher in Tropics and lower over Southern Ocean; Generally a lower frequency over mid-latitude Land

How much of the observed difference is random?

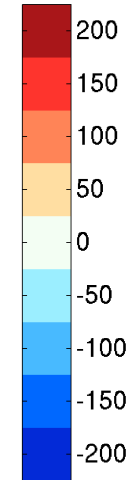
High clouds < 440 hPa

MODIS C5 minus C6 avg difference Aug 2012 - Daytime high clouds (9.63 hPa)



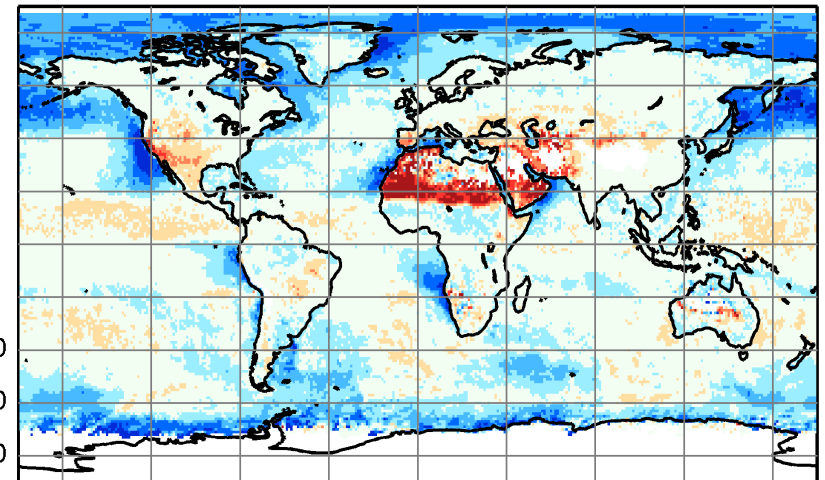
1-degree

[hPa]



Low clouds > 680 hPa

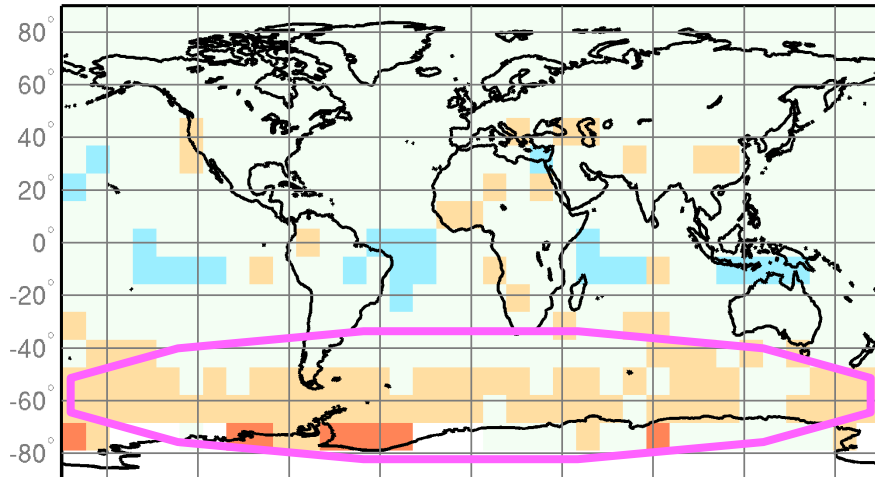
MODIS C5 minus C6 avg difference Aug 2012 - Daytime low clouds (-13.44 hPa)



Increase grid resolution by an order of magnitude to remove random differences

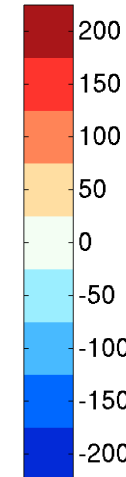
High clouds < 440 hPa

MODIS C5 minus C6 avg difference Aug 2012 - Daytime high clouds (5.62 hPa)



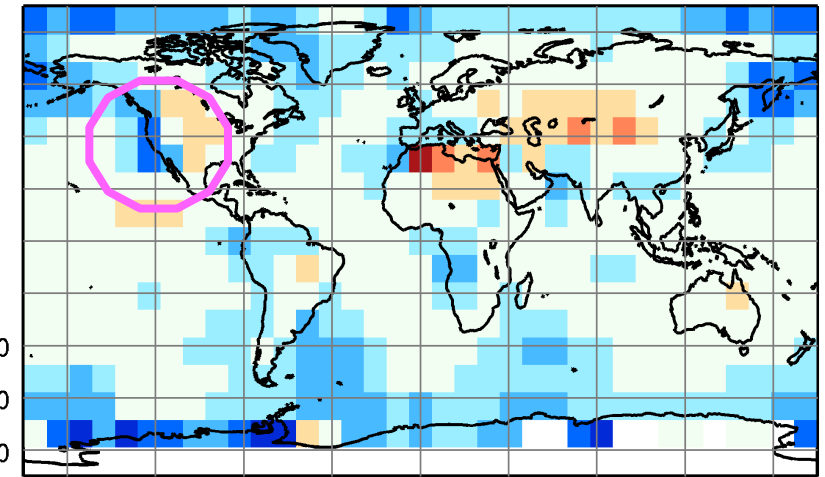
10-degree

[hPa]



Low clouds > 680 hPa

MODIS C5 minus C6 avg difference Aug 2012 - Daytime low clouds (-16.67 hPa)



How seasonal is systematic differences?

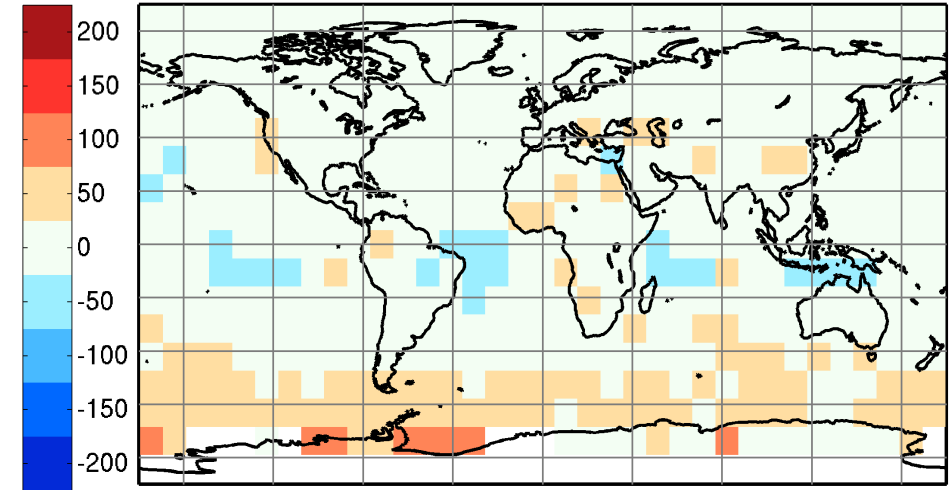
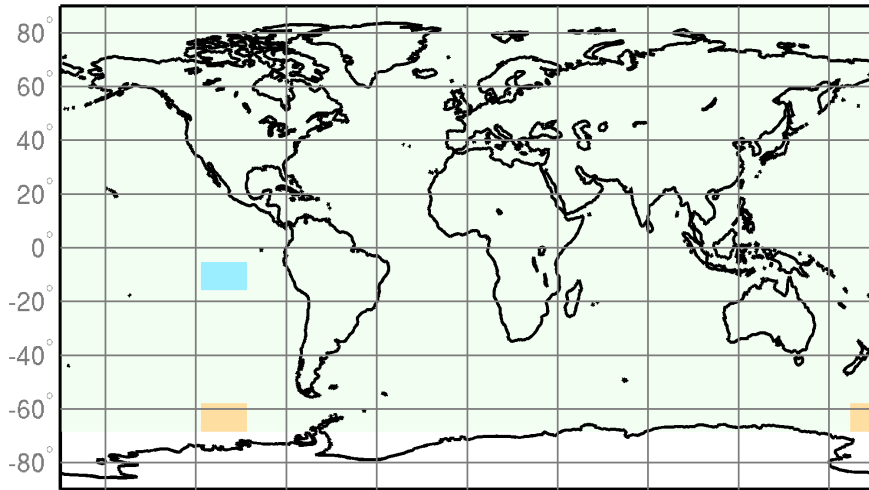
High clouds < 440 hPa
C5 CTP monthly mean

Spring (August 2013)

Summer (August 2013)

MODIS C5 minus C6 avg difference May 2012 - Daytime high clouds (1.95 hPa)

[hPa] MODIS C5 minus C6 avg difference Aug 2012 - Daytime high clouds (5.62 hPa)

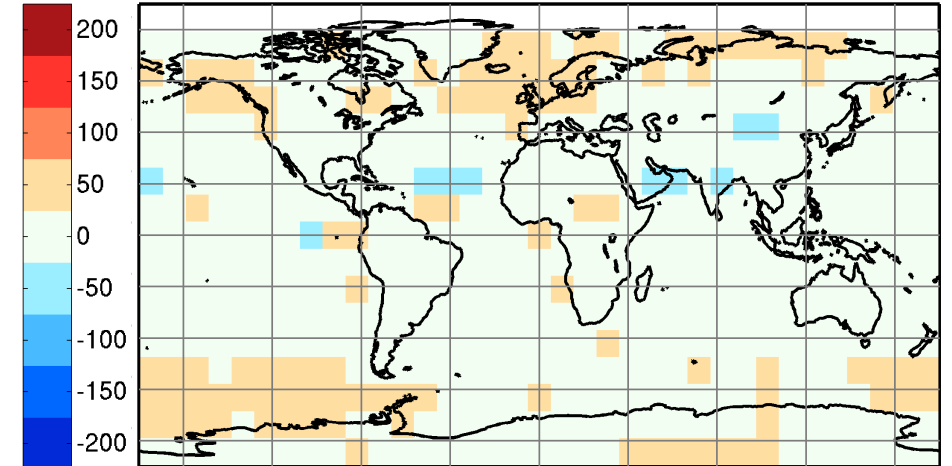
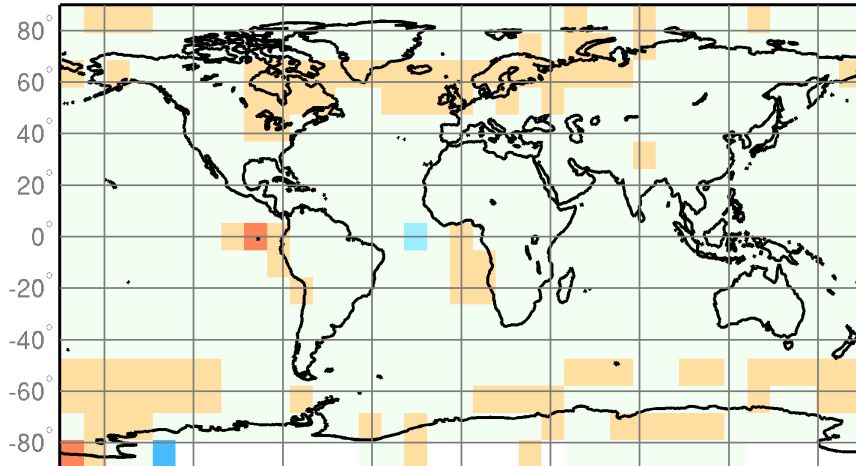


Autumn (November 2012)

Winter (January 2013)

MODIS C5 minus C6 avg difference Nov 2012 - Daytime high clouds (9.76 hPa)

[hPa] MODIS C5 minus C6 avg difference Jan 2013 - Daytime high clouds (6.36 hPa)



Analyzing Systematic Differences

Instrument:

AIRS + IASI + CrIS

Retrieval Algorithm:

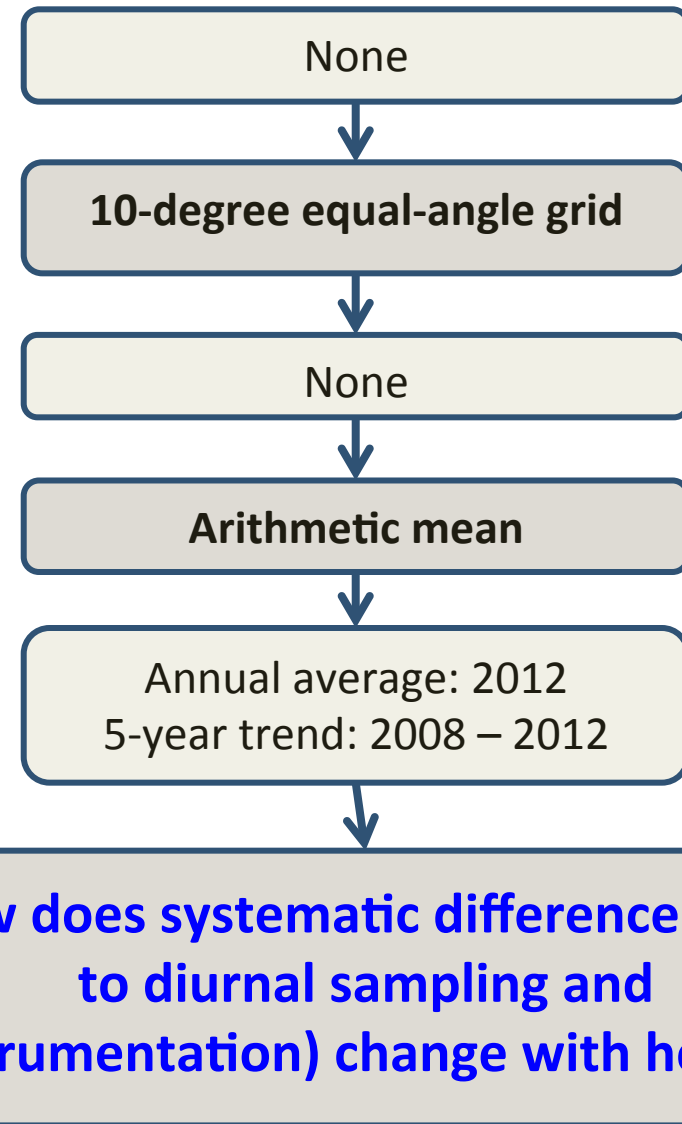
UW Dual-Regression

Parameter:

Temperature Profile

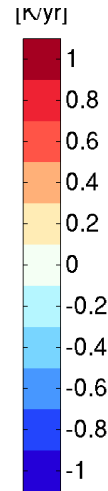
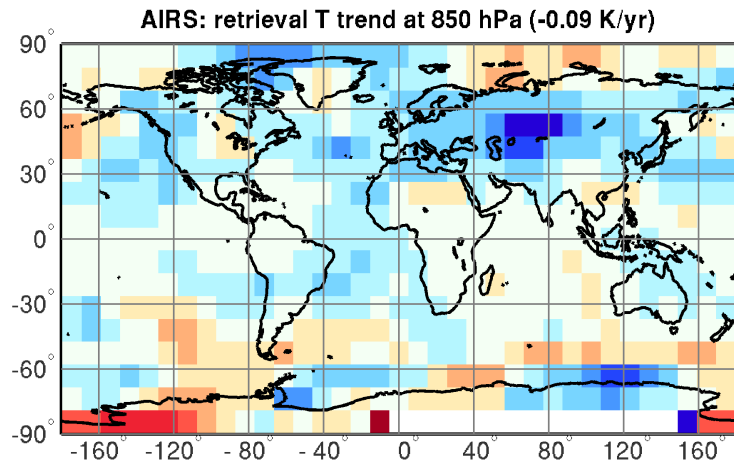
[K]

Using model fields
(from GDAS)
as reference

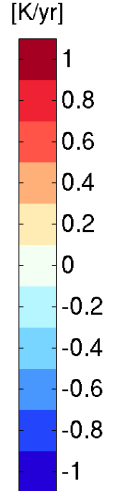
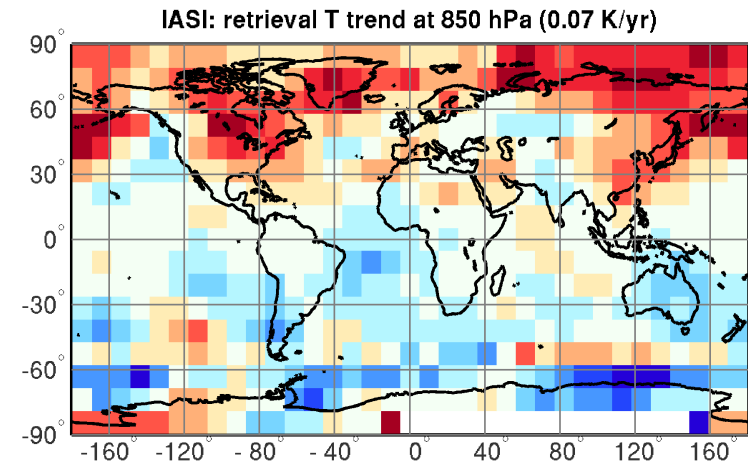


5-yr Trend (2008–2012) at 850 hPa (K/yr)

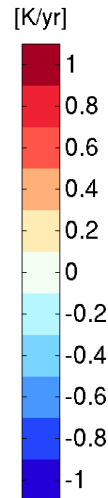
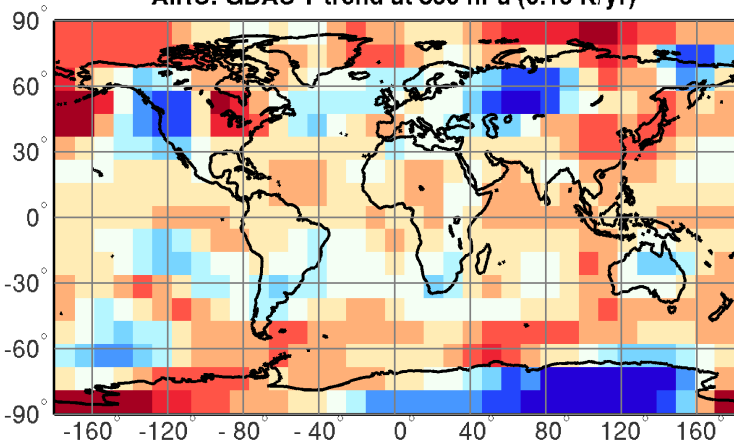
AIRS @ 13:30 LST



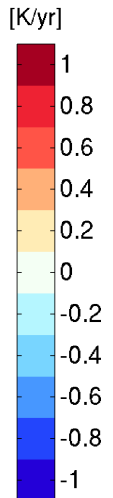
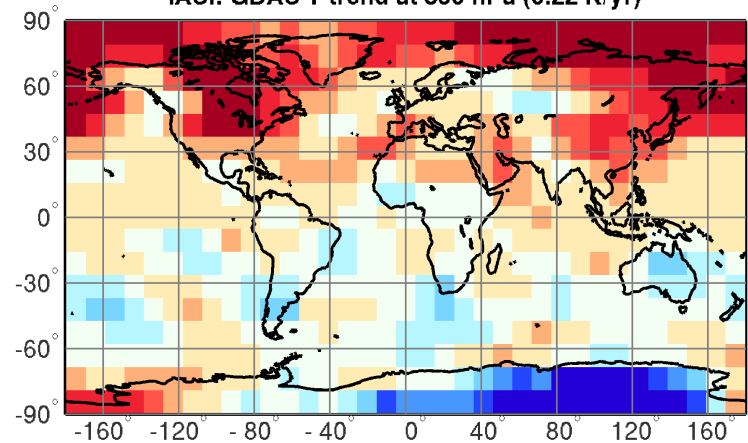
IASI @ 09:30 LST



AIRS: GDAS T trend at 850 hPa (0.15 K/yr)



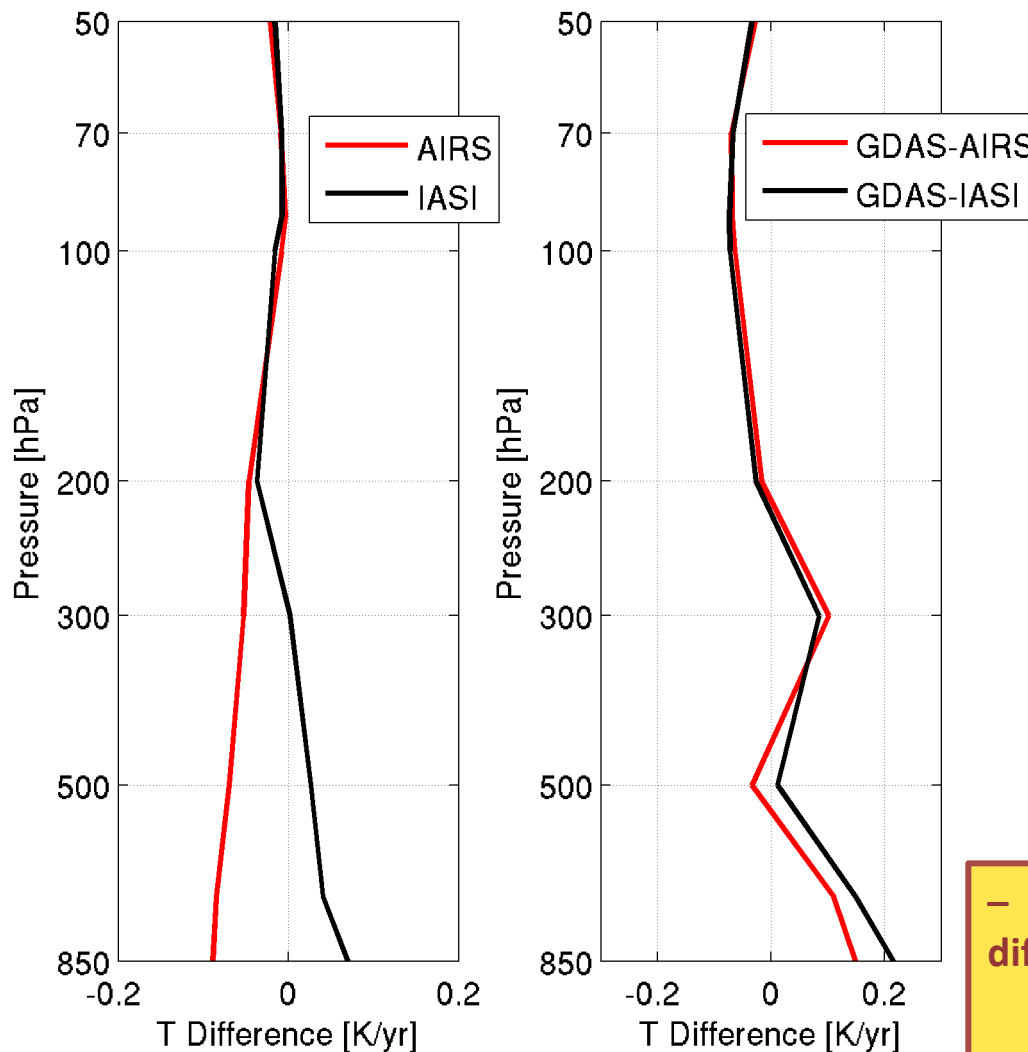
IASI: GDAS T trend at 850 hPa (0.22 K/yr)



IASI-AIRS Trend difference is 0.16 K/yr while GDAS Trend difference is 0.07 K/yr

At 850 hPa half of the observed difference between AIRS & IASI is due to diurnal sampling differences

Analyzing Systematic Differences



How does diurnal sampling differences affect product comparisons at different heights?

Global weighted (by cosine of the latitude) mean along a pressure profile.

Isolating two sources of systematic differences by comparison with model data (as reference)

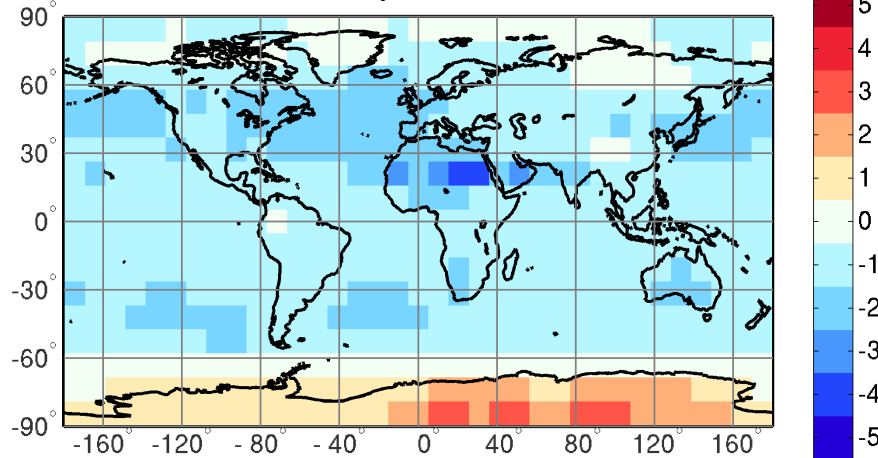
- (1) Diurnal sampling
- (2) instrument

- Uncertainty due to diurnal sampling differences affects tropospheric retrievals only
- Possible implications for evaluating/analyzing multi-instrument low-cloud retrievals

Systematic difference caused by instrument alone

AIRS @ 13:30

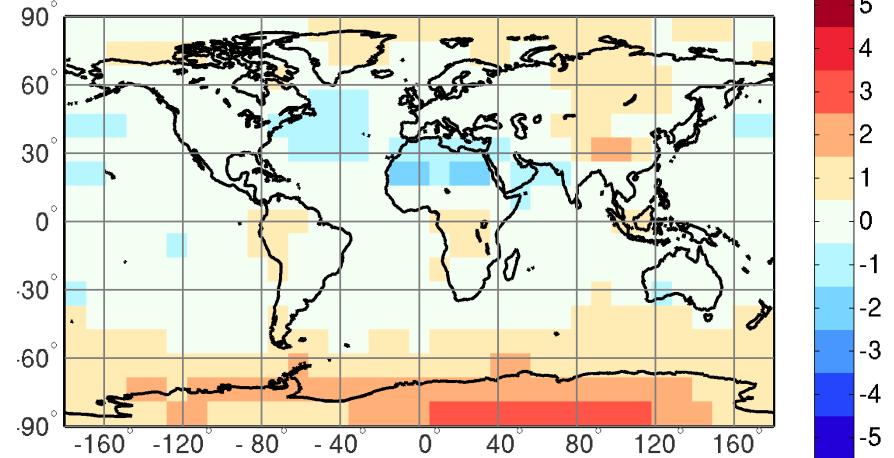
AIRS-GDAS daily difference: -1.05 K



300 hPa

CrIS @ 13:30

CrIS-GDAS daily difference: 0.08 K



Annual mean of daily differences

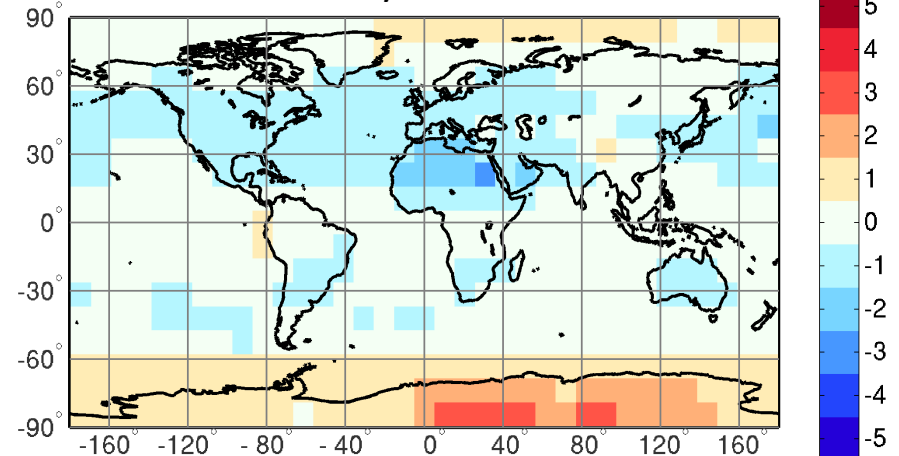
Daily difference = Instrument minus Space-Time-Interpolated Model Field (diurnal sampling effect removed)

AIRS is a grating spectrometer

IASI + CrIS are Michelson Interferometers

IASI @ 09:30

IASI-GDAS daily difference: -0.25 K

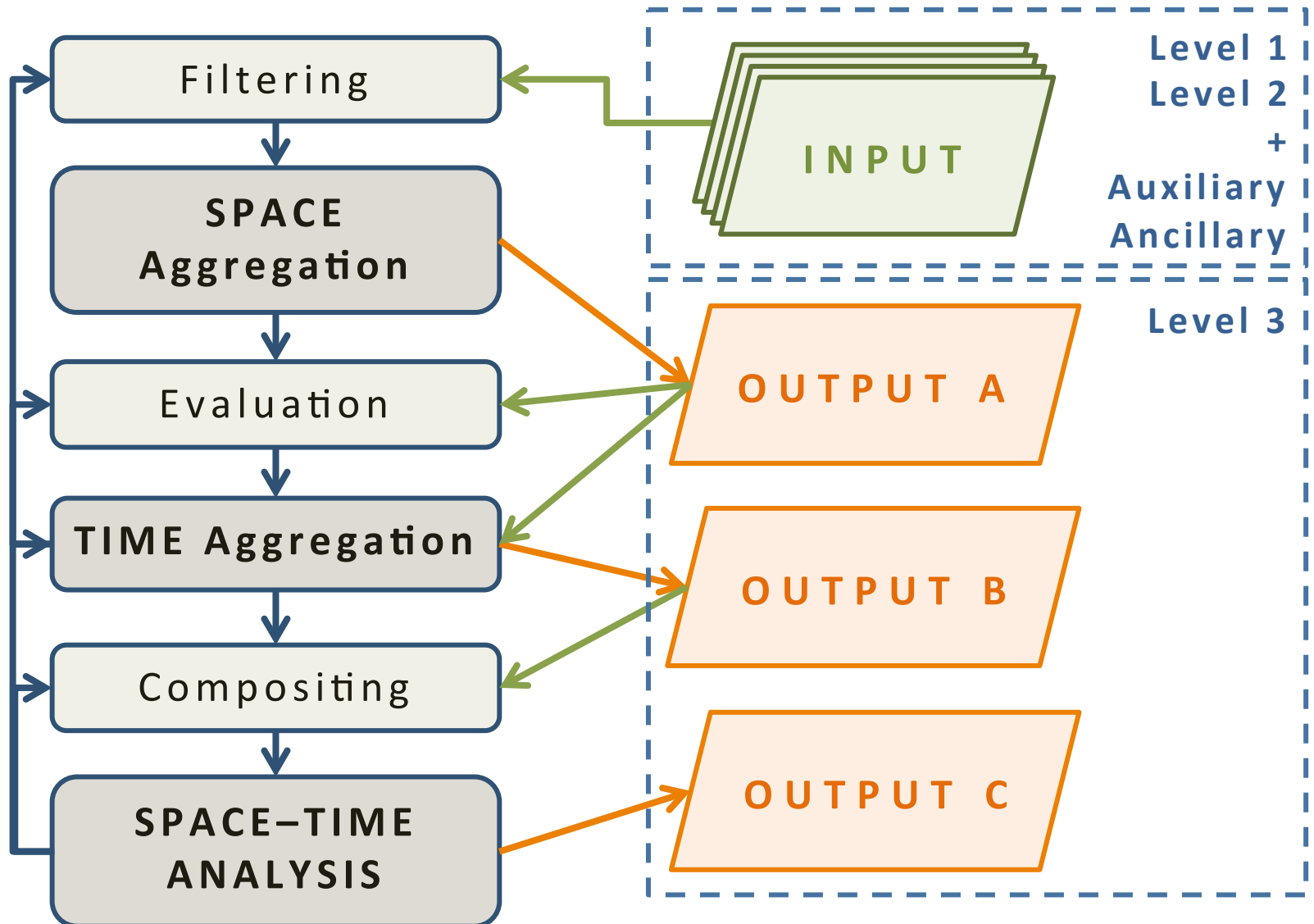


IASI and CrIS have stronger correlation with each other than with AIRS

Uncertainty caused by L-3 methods

UW Space–Time–Gridding (STG) Framework

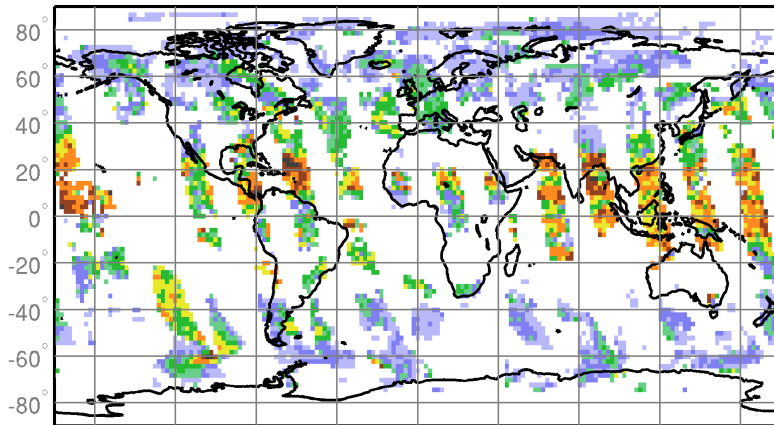
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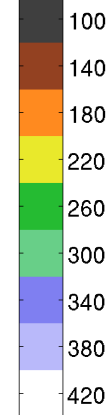
Propagation of L3 Uncertainty: Spatial Scale

MODIS C5: Daily (3 Sep 2012) high CTP difference on different resolution grids

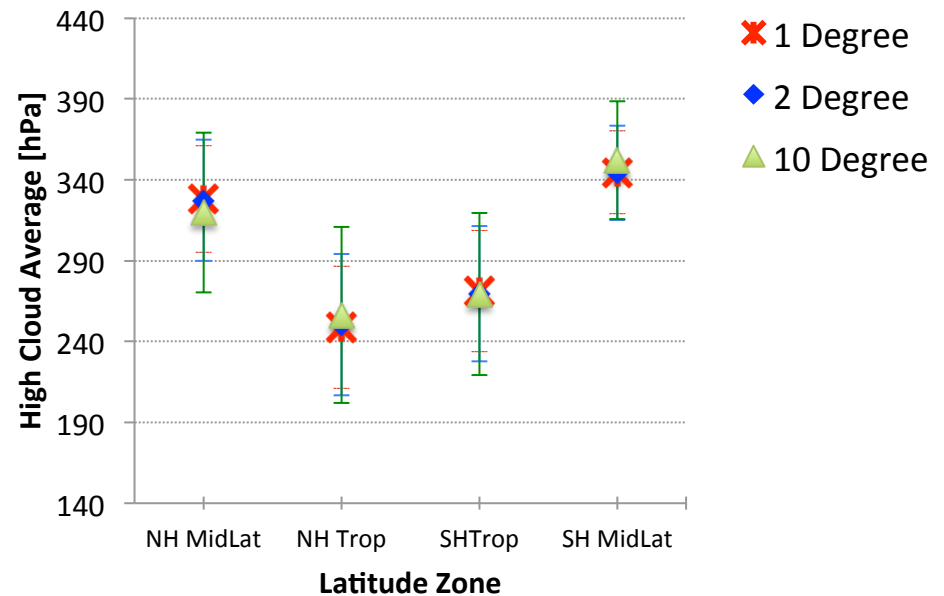
Daily average of high CTP on 2° grid



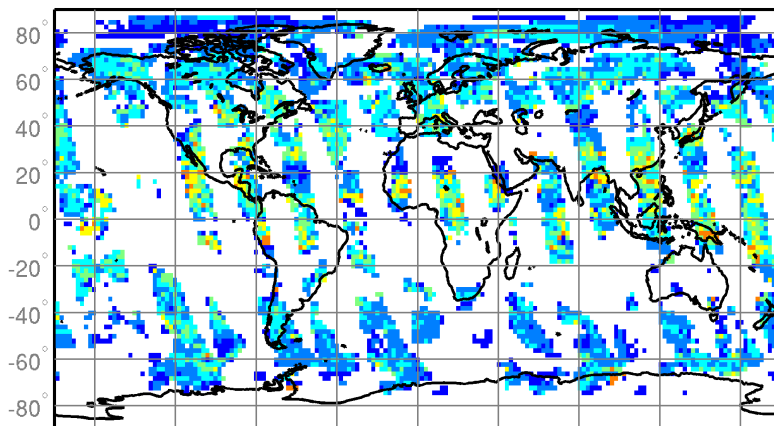
[hPa]



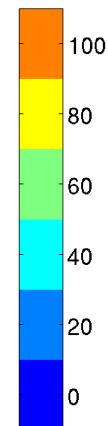
How critical is grid-cell size to global/zonal statistics?



Daily uncertainty (SD) of high CTP on 2° grid



[hPa]



Mean = mean of grid cell means

Error bars = mean of grid cell uncertainties (SD)

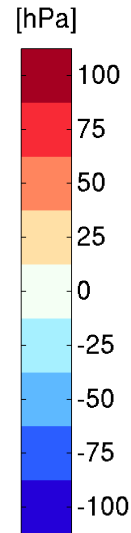
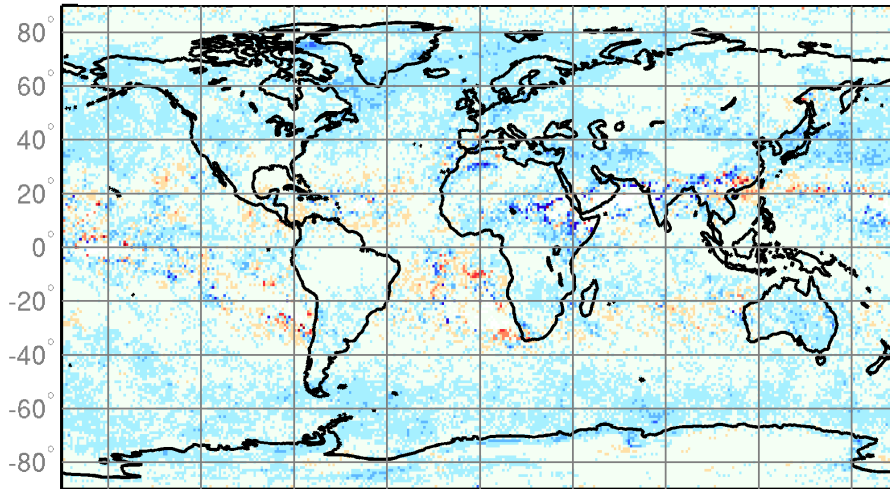
For zonal statistics of CTP, spatial resolution of L3 product (daily space-time aggregate) is not critical

Note: a larger spatial scale will affect the rate at which the probability in each cell changes, i.e. a 10-degree (1-degree) cell will require a higher (lower) degree of change across space before a change in the daily estimate (mean of pdf) is observed

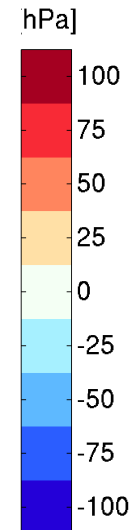
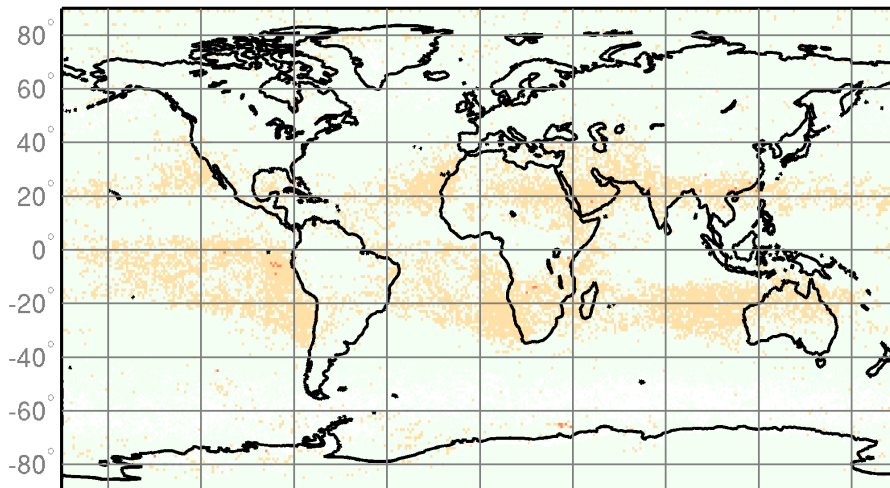
Propagation of L3 Uncertainty: Temporal Scale

AIRS DR: Monthly high CTP difference on 1-degree grid

Feb 2012: difference of two types of monthly means



2003–2012: 10-year mean difference of two monthly means



How does L3 aggregation affect product comparisons?

How much of the L3 product differences smooth out over time? (i.e., how much of the L3 product differences are random?)

Type 1 monthly mean = monthly mean of daily means

Type 2 monthly mean = weighted (by % cloudiness) monthly mean of daily means

Difference = Type 1 minus Type 2

Uncertainty due to systematic effects becomes evident once temporal scale is increased

Propagation of L3 Uncertainty: Temporal Scale

AIRS DR Temperature @ 300 hPa difference on 10-degree grid

What constitutes a monthly mean?

Does it matter which combination of days we use?

Type 1 = monthly mean of 15 even-numbered days

Type 2 = monthly mean of 15 odd-numbered days

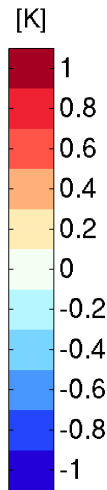
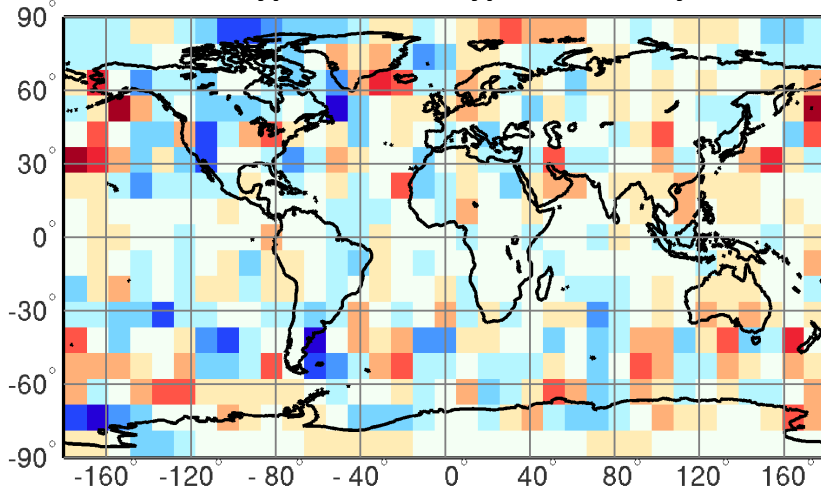
Type 3 = monthly mean of first 15 days

Type 4 = monthly mean of last 15 days

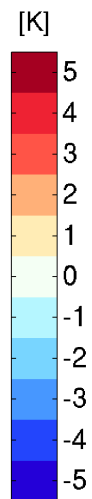
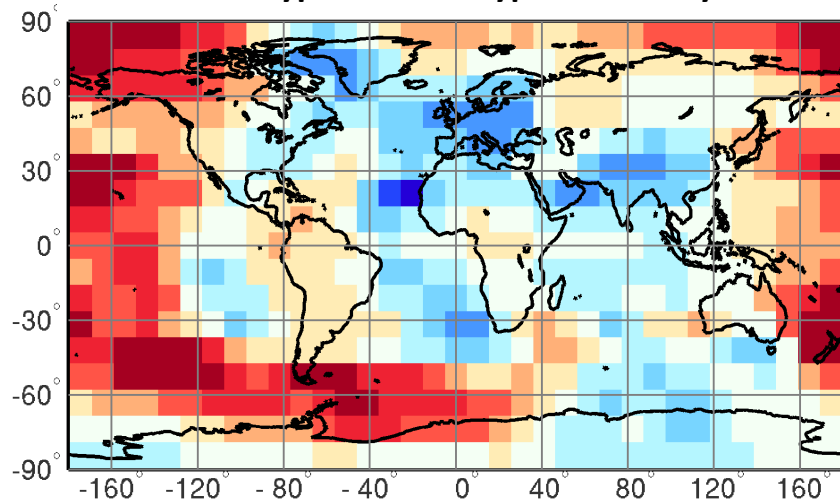
A Monthly mean is sensitive to the combination of daily means.

An even distribution of days across a month (irrespective of exact matching) results in **random differences** only. Two different cluster distribution of days across a month can result in **systematic differences**.

Feb 2012: Type 1 minus Type 2 monthly means

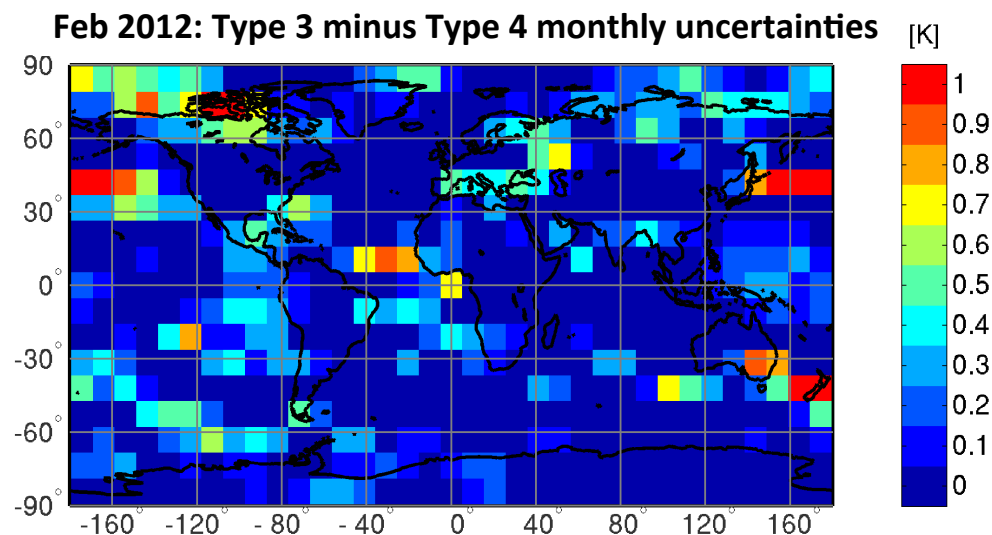
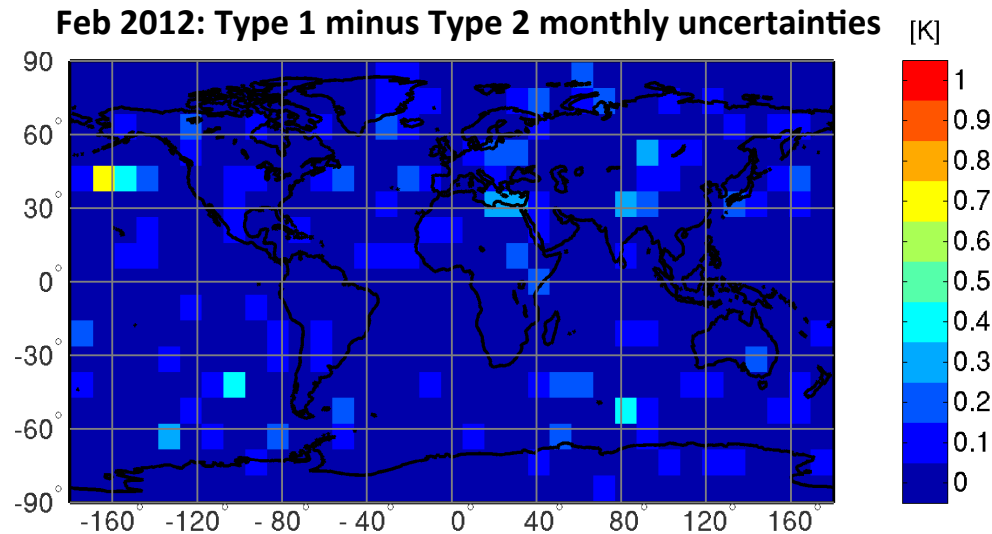


Feb 2012: Type 3 minus Type 4 monthly means



Propagation of L3 Uncertainty: Temporal Scale

AIRS DR Temperature @ 300 hPa difference on 10-degree grid



What causes these observed differences in monthly means?

How can daily L-3 uncertainties help clarify differences?

Average monthly uncertainty = Average of daily SD per grid cell

A high average uncertainty indicates a high variability in parameter state (due to weather events). If a time-aggregate sample of daily averages is not evenly distributed, but clustered around weather events, L-3 uncertainty can no longer be attributed to random effects alone.

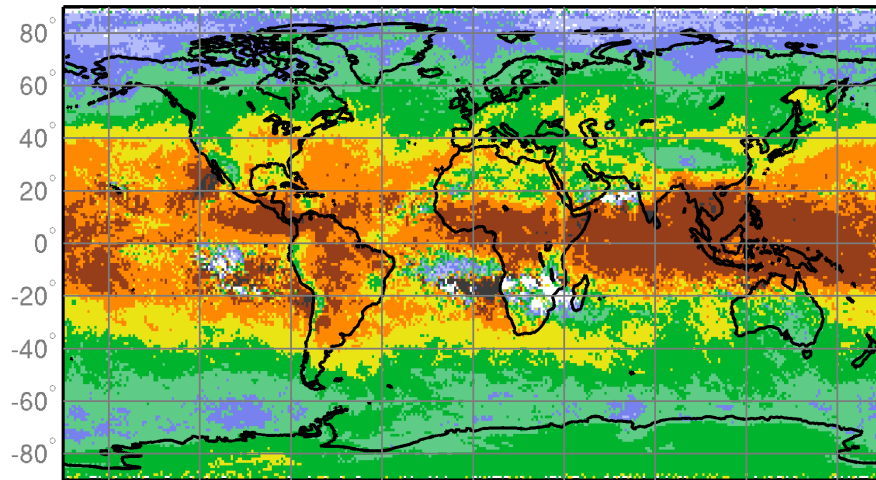
L-3 uncertainty metrics can help design aggregation strategies

Multi-source Multi-instrument Global comparisons

Multi-Instrument Global Comparisons

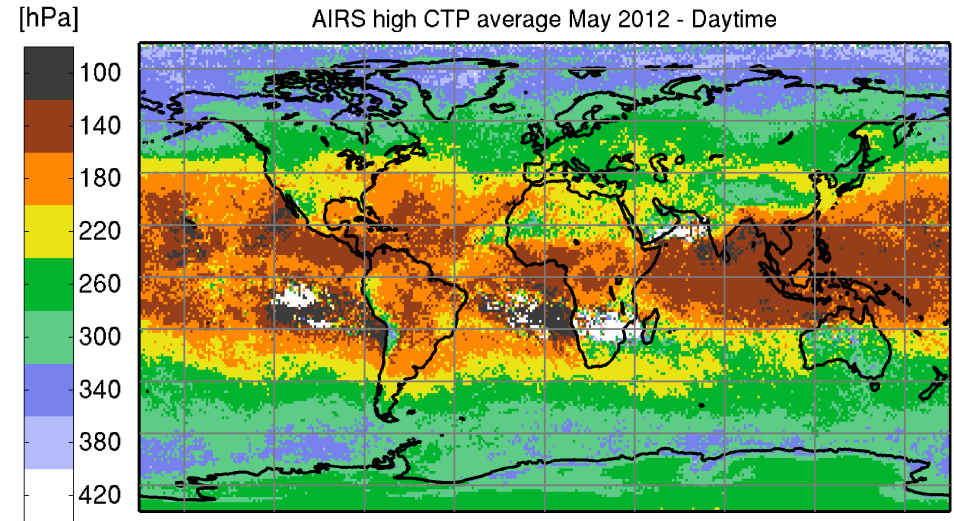
Sounder (IASI) @ 09:30 (UW DR)

IASI high CTP average May 2012 - Daytime



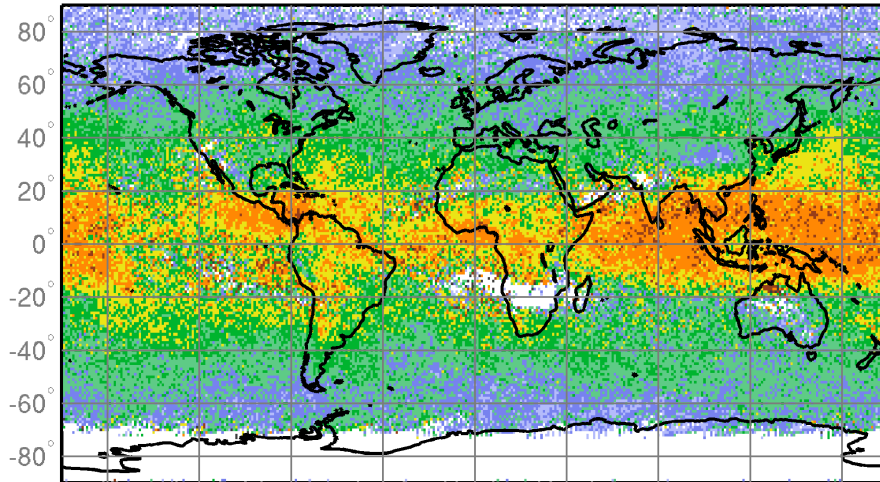
Sounder (AIRS) @ 13:30 (UW DR)

AIRS high CTP average May 2012 - Daytime



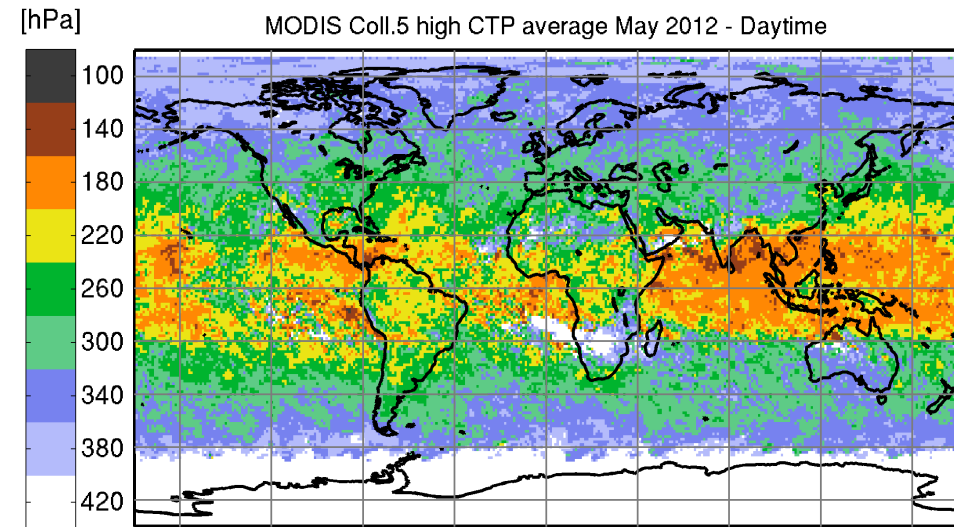
Imager (AVHRR) @ 09:30 (Optimal Estimation)

AVHRR high CTP average May 2012 - day

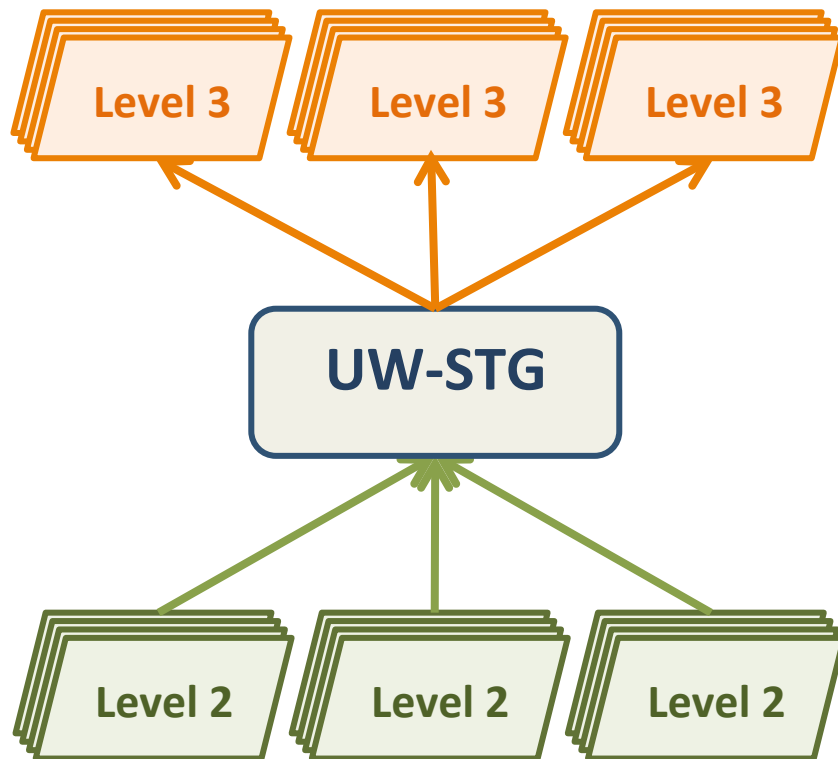


Imager (MODIS) @ 13:30 (CO₂ slicing)

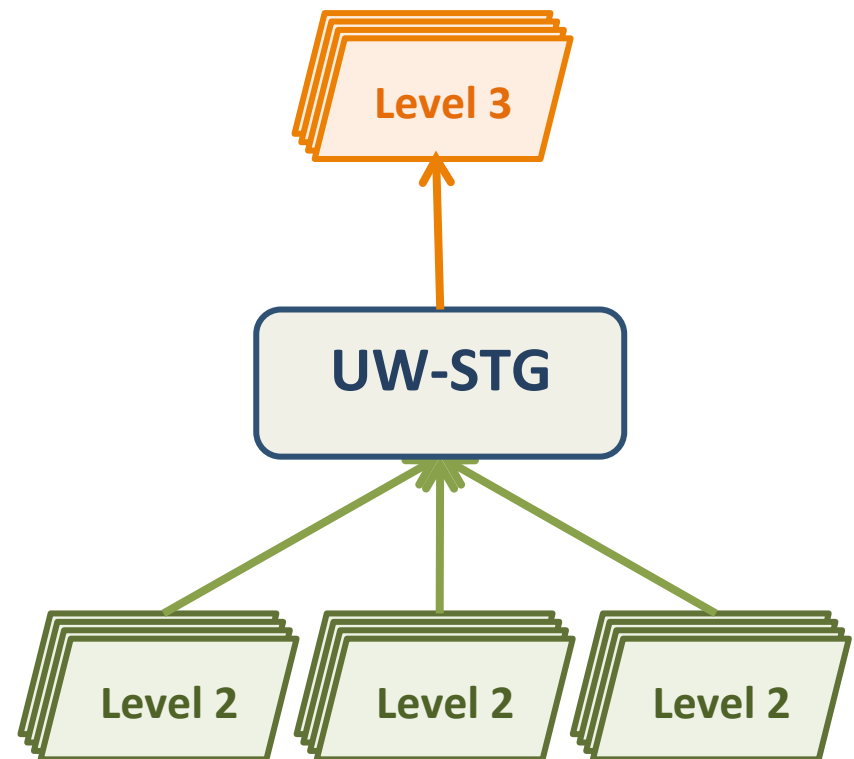
MODIS Coll.5 high CTP average May 2012 - Daytime



STG supports different approaches to production of aggregated products



Simple products from single sources



Composite product with (weighted) input from multiple sources

SUMMARY and FUTURE WORK

- Uncertainty can be traced, characterized, understood and managed. It can be used to help design aggregation algorithms.
 - STG is transparent, iterative, fast, independent of instrument and dynamic (allows implementation of user-defined aggregation strategies).
 - A modular approach to aggregation facilitates a systematic and traceable investigation into the behavior of data and geophysical properties over space-time
 - Aggregation can be used to
 - Better understand L2 differences/algorithms/uncertainties
 - Make user-specific L3 products with high information content and good uncertainty estimates
 - Aggregation is not limited to the averaging of L2 pixels. L3 can be higher level information products that draw on data from multiple parameters/sources)
-
- Continue to characterize sources of uncertainty
 - Investigate methods for working with multiple observations at different times.
 - Evaluate cloud parameter differences between Imager-Sounder pairs (MODIS/AIRS, AVHRR/IASI, VIIRS/CrIS). Develop composite imager-sounder cloud products
 - Build long-term sounder (HIRS/AIRS/CrIS/IASI) cloud record
 - Investigate the effect of ancillary data resolution on L-3 products